

Developing Physics-oriented Diagnostic Tools for Model Evaluation and Improvement

PI: Zhuo Wang,
Department of Atmospheric Sciences
University of Illinois at Urbana-Champaign
Co-PIs: Stan Benjamin (NOAA ESRL)
Melinda Peng (NRL Monterey)
Min Zhao (NOAA GFDL)

Model Evaluation

- Evaluation of model forecasts is an indispensable component of the model development effort.
- Standard performance-oriented metrics, such as ACC and RMSE, are routinely used in operational centers for forecast verification. NCAR DTC also developed model evaluation tools (MET).
- Physics-oriented diagnostics focus on the critical processes or phenomena and, through evaluation of key variables, shed light on the deficiencies of the physical parameterization or other errors in a model.
- In brevity, physics-oriented evaluation not only provides information on how well a model performs but also on why a model may fail in a certain aspect.
- There is an increasing demand in recent years in the community to develop and adopt physics-oriented diagnostics

Over-arching Objective

- **Objective:** develop physics-oriented diagnostic tools to assist the development of the NOAA's Next Generation Global Prediction System (NGGPS) under the R2O Initiative
- **Expected outcome:** a suite of diagnostic tools with general applicability across models
- In-depth evaluation of model forecasts using physics-oriented metrics helps to “*improve the physical parameterizations to allow for efficient, accurate and more complete representations of physical processes and their interactions across scales*” and helps to “*achieve a world class global prediction system*”, which is highly relevant to the priority of the R2O initiative.

Outline of Physics-based Evaluations:

Two levels of products

1. Level 1: Prominent Motion Systems

- Tropical Cyclones (**ongoing**)
- MJO (**ongoing**)
- Blocking (**to be developed**)
- Teleconnection (**to be developed**)

2. Level 2: Specific Physical Processes

- Cumulus parameterization: CWV-precip relationship, Q1/Q2 diagnoses (**ongoing**)
- Model representation of different cloud regimes (**to be developed**)

Two Phases

- Phase 1: The products will be developed and tested using the GEFS reforecasts.
- Phase 2: Products will be generalized to different models; an inter-model comparison will be carried out by taking advantage of the HIWPP project

Data: Global Ensemble Forecasting System Reforecast-2 (GEFS-R)

- a) GEFS-R uses **a fixed version** (9.0.1) of the GEFS → useful information on the operational GEFS;
- b) With the forecast lead up to 16 days
- c) A multi-decadal reforecast database (**~ 30 years**) with larger sample size better evaluates model performance and predictability,
- d) Ensemble forecasts make probabilistic forecast possible.

Data: GEFS-R

Dynamical model	Resolution	Ensembles
<i>Week 1 (Day 0 - Day 7)</i>	<i>T254L42 (~ 40 km) 42 levels in vertical</i>	<i>1 control + 10 perturbed members</i>
<i>Week 2 (Day 8 - Day 16)</i>	<i>T190L42 (~ 54 km) 42 levels in vertical</i>	<i>1 control + 10 perturbed members</i>
<i>SST analysis: assimilates observations from past 7 days SST forecast: damped with an e-folding time of 90 days</i>		

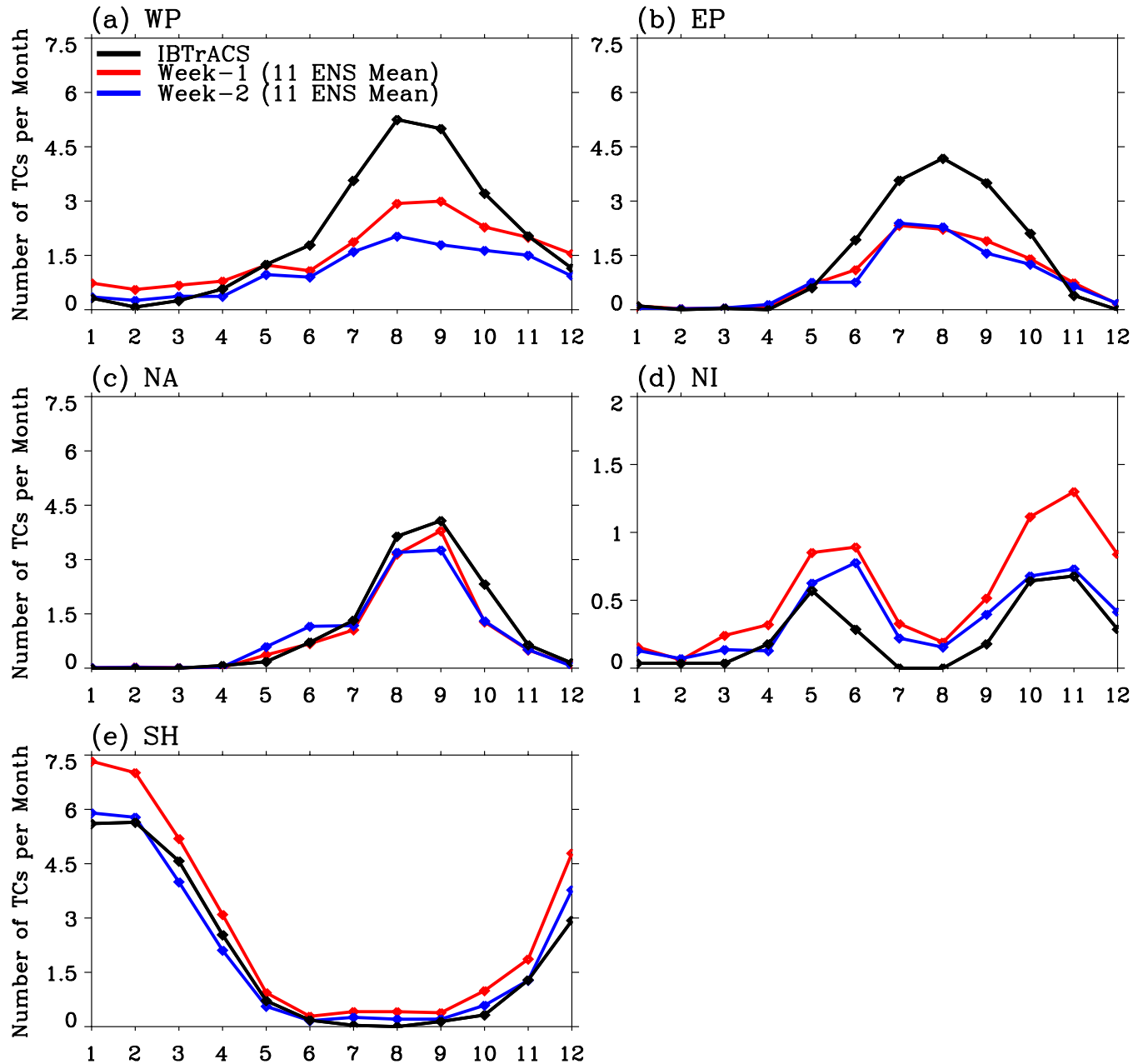
Time Period	<i>01 Jan – 31 Dec (1985 – 2012)</i>
Observation/Reanalysis	<i>IBTrACS , ERA-interim CMORPH, SSML, etc.</i>
Region	<i>Global (5 basins)</i>

1.1 Tropical Cyclones

Detection of Model TCs

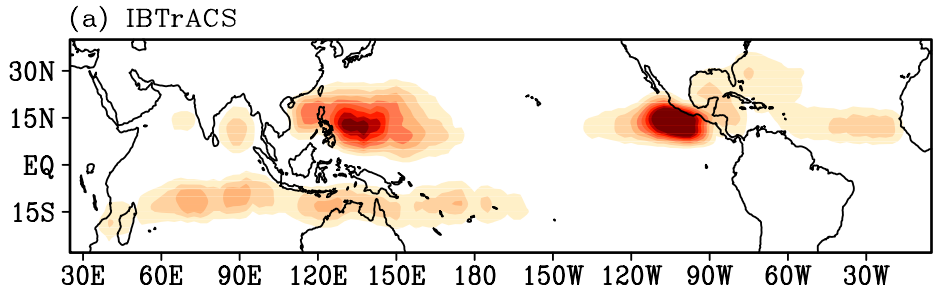
- **GFDL vortex tracker** (*Marchok 2002*): 6-hourly TC tracks
- **Exclude:** i) short-lived TC (≤ 48 h), ii) TC with maximum intensity $\leq 16.5 \text{ m s}^{-1}$ (~ 32 knots), iii) genesis latitude higher than 40° , iv) warm core lasting ≥ 48 h during its lifetime
- Observation: IBTrACS data, 1985 - 2012
- Categorization of model TCs: hit, early genesis (EG), late genesis (LG), and false alarms

Seasonality of TC Genesis in Diff. Basins

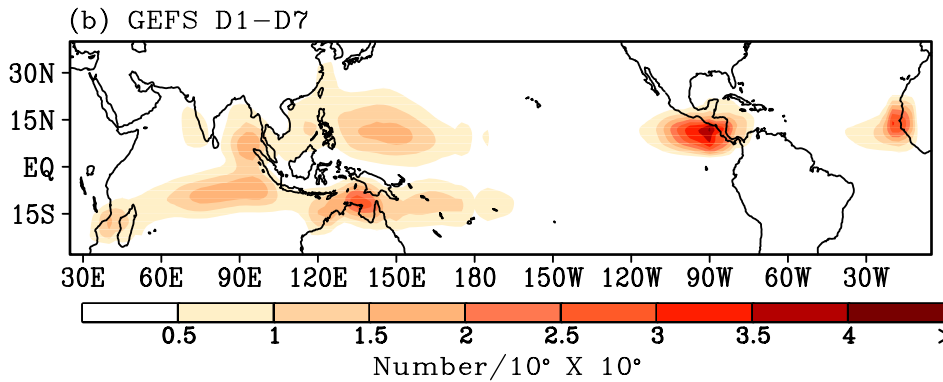


Climatological Mean TC Genesis Density (1985–2012)

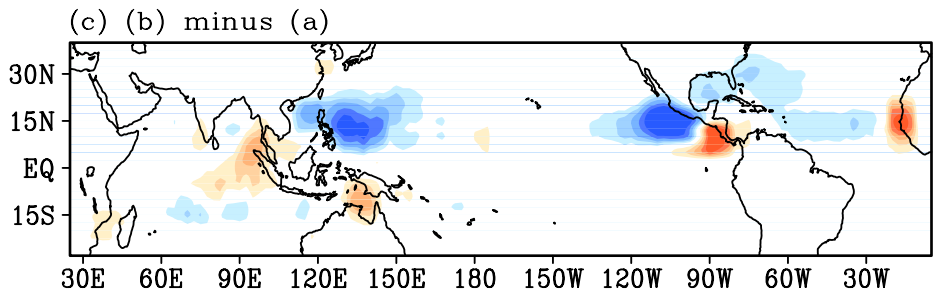
IBTrACS



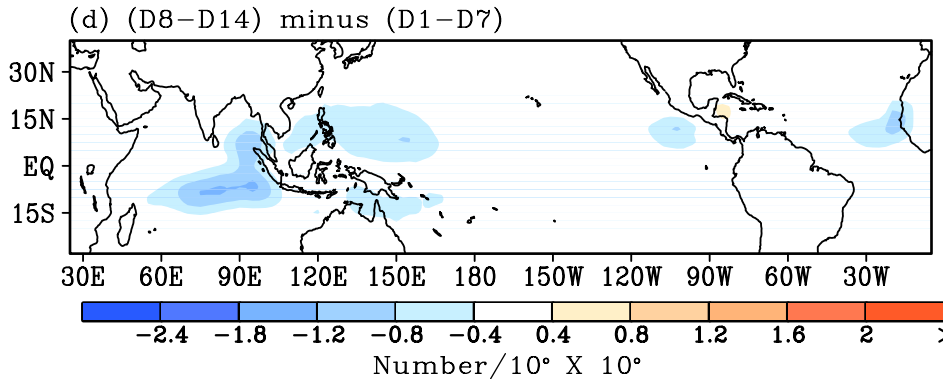
Week-1



“Week-1” - Obv



“Wk-2” - “Wk-1”



Genesis density function: 2D distribution of genesis frequency

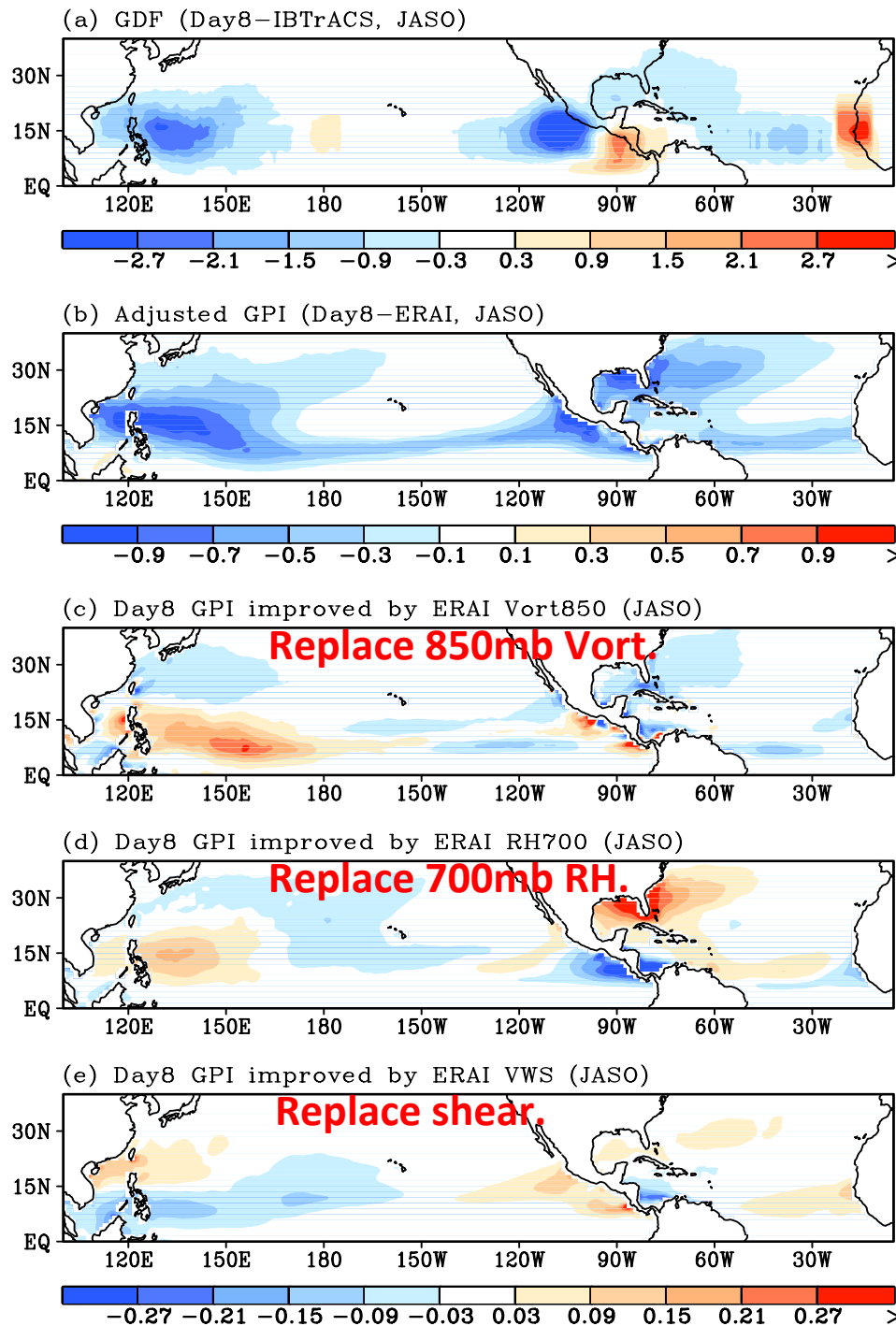
The TCG distribution in GEFS is broadly consistent with IBTrACS, but quantitative diff. are also evident.

Unit:
of TCs forming within a 10-degree box per year

Genesis Potential Index

(Emanuel and Nolan 2004)

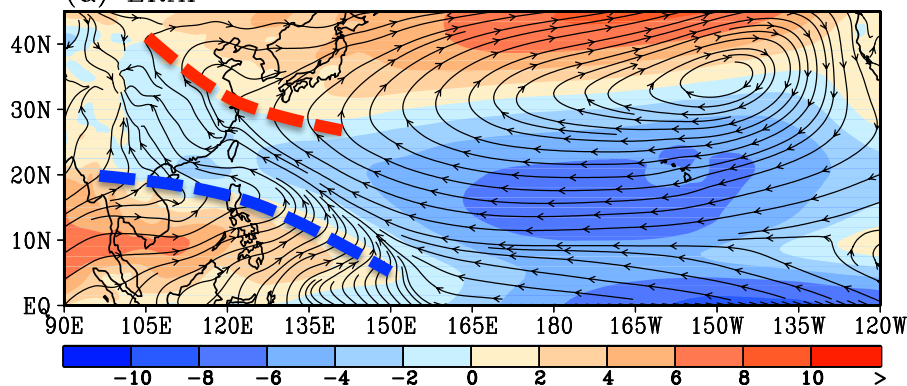
$$GP = |10^5 \eta|^{\frac{3}{2}} \left(\frac{RH}{50}\right)^3 \left(\frac{PI}{70}\right)^3 (1 + 0.1VWS)^{-2}$$



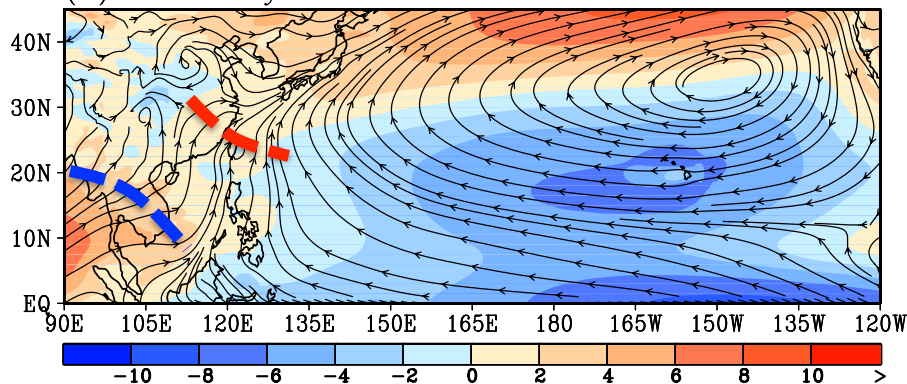
- 850 hPa vorticity, 700 hPa RH, VWS and PI.
- GPI is a function of the environmental conditions, can be used as a proxy for genesis probability.
- Comparison of GPI derived from GFS-R and ERA1: How much can we attribute the errors in TCG to errors in the large-scale environment?

Long-Term Mean Wind (850 mb, Jul - Oct)

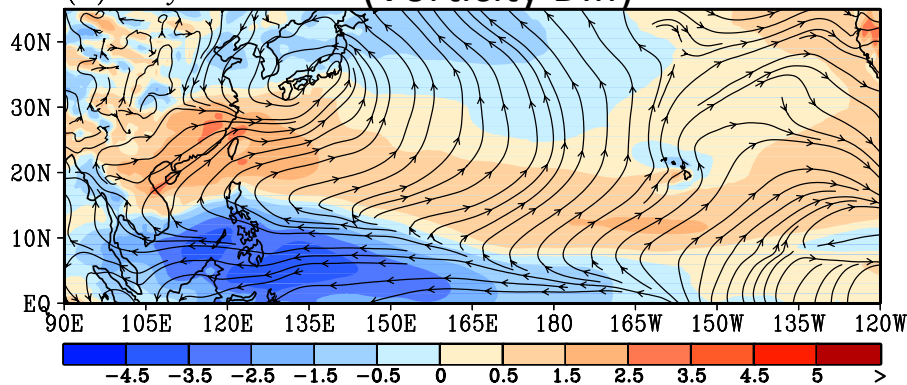
(a) ERAI



(b) GEFS Day8 Fcst



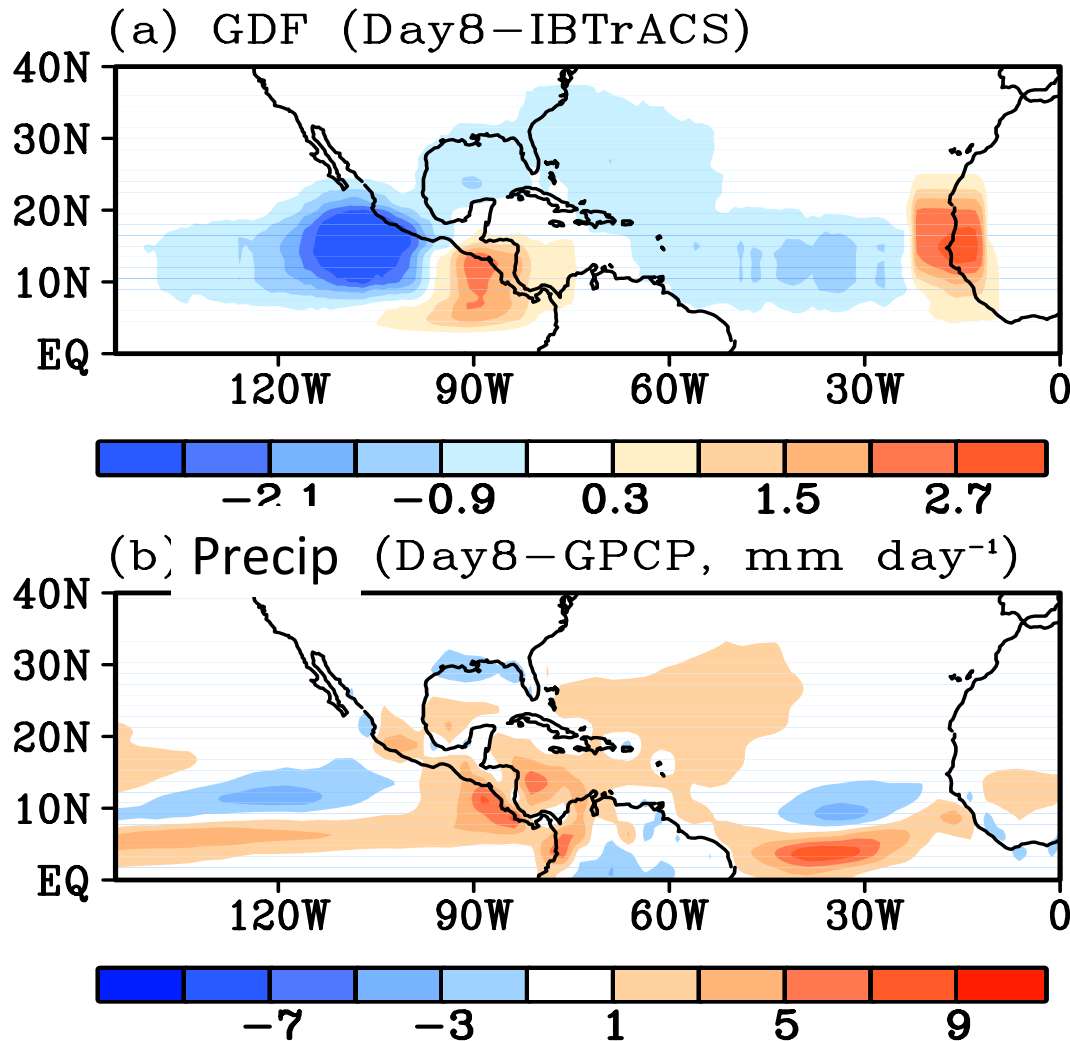
(c) Day8 - ERAI (Vorticity Diff)



850 hPa Wind

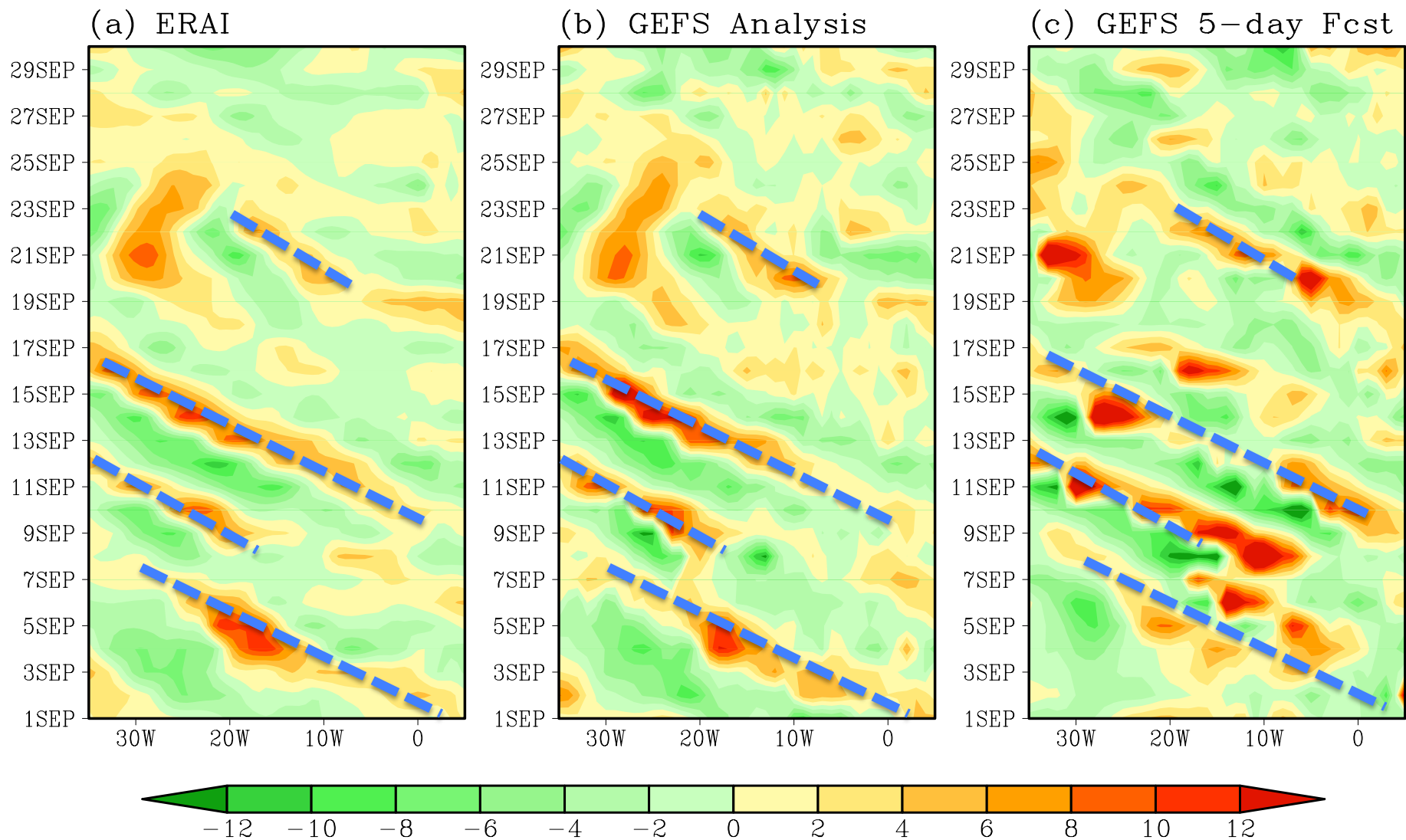
- The monsoon trough is too weak in GEFS, leading to a negative bias of relative vorticity and drier midlevel atmosphere over the western north Pacific.
- The weaker subtropical ridge near Japan contributes to a positive bias in extratropical cyclogenesis in that region (not shown).
- An improved mean state can improve TC forecasting.

East Pacific and Atlantic

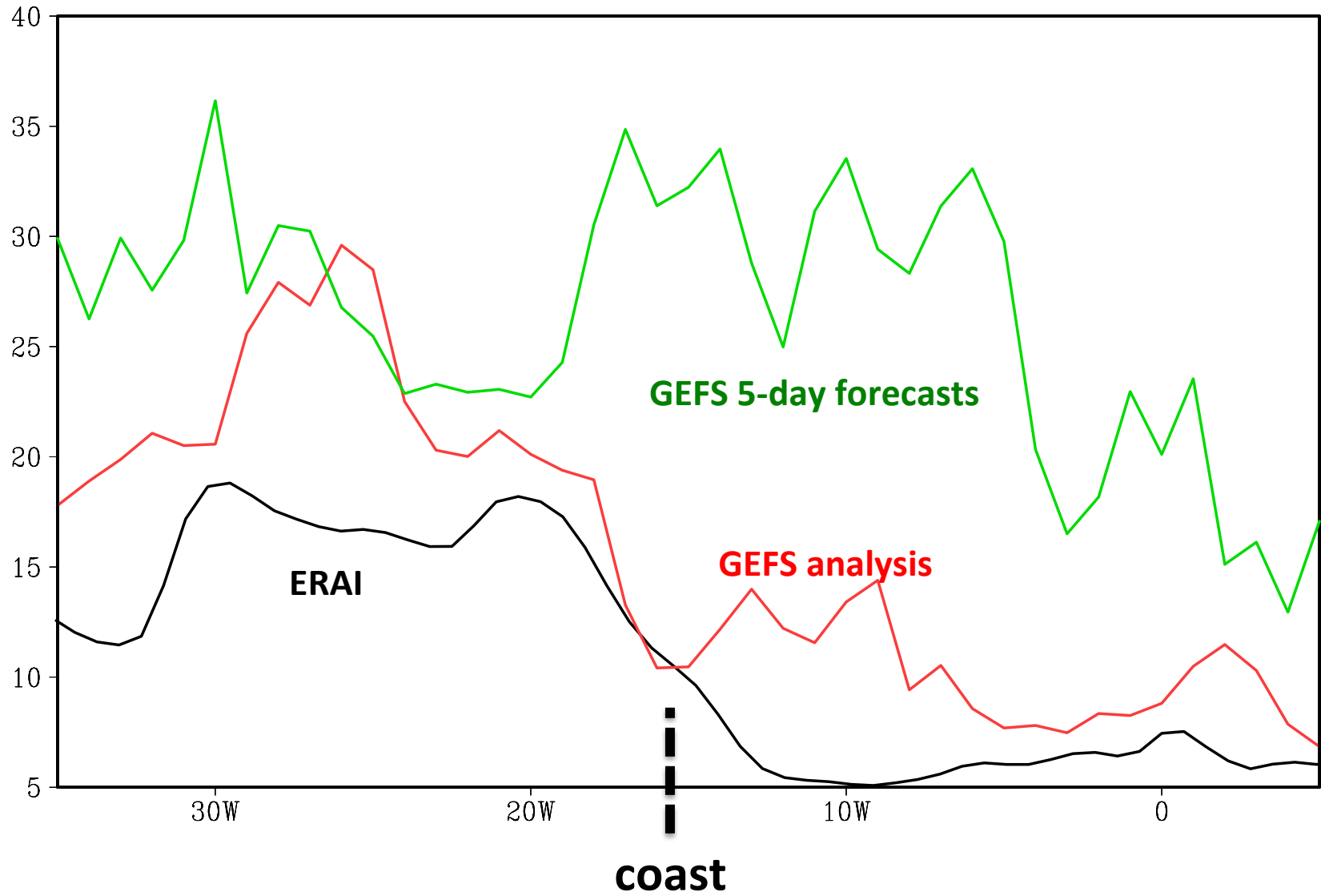


- The ITCZ is displaced southward over the E. Pac and the Atlantic.
- The tip of the ITCZ is most unstable for ITCZ breakdown.
- Precipitation is overpredicted over West Sahel.

Hovmoller Diagram of V850 (15N, 2010)

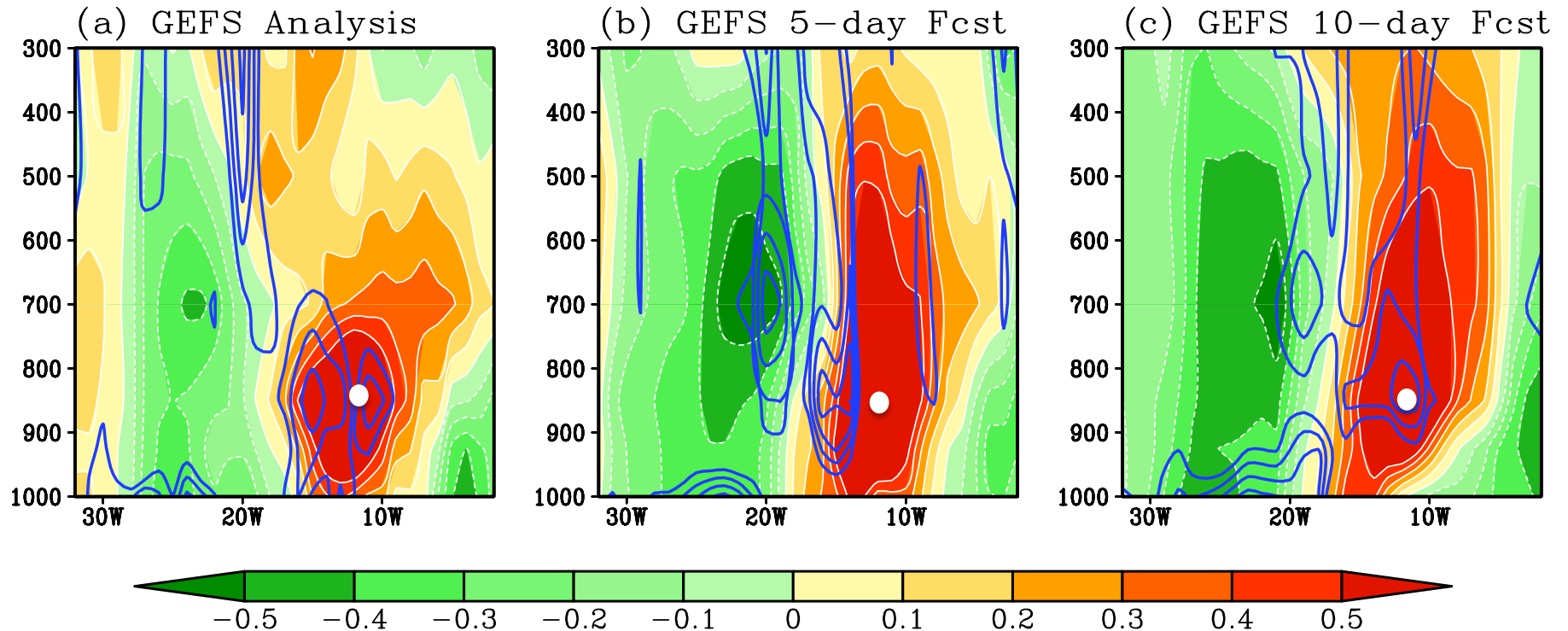


Variance of V850 along 15N (summer 2010)



Vertical Cross Section of Correlation (summer 2010)

Corr. of V550 (01AUG2010–30SEP2010, 15N)



One-point correlation of V, with the reference point: 850 hPa, (15N, 12W)

Hyperactive convection over land leads to stronger AEWs of deeper vertical structure, and the early TC development (false alarms) off the coast.

1.2 MJO: boreal summer

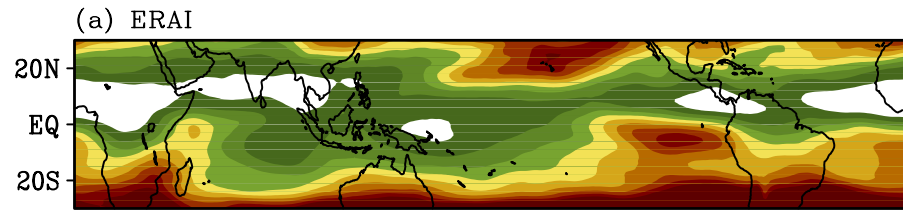
Why are we interested in the MJO?

- The dominant mode of intraseasonal variability in the tropics, and plays a significant role in modulating the tropical weather and climate:
 - strong impacts on Asian and Australian monsoon and moderate impacts on the North American and South American monsoon (Lau and Chan 1986; Kiladis and Weickmann 1992; Mo 2000; Higgins et al. 2000)
 - plays an active role in the onset and development of ENSO (e.g., Zhang and Gottschalck 2002)
 - modulates tropical convection and tropical cyclone activity over the East Pacific and the Atlantic
 - has remote impacts in the extratropics and affects the midlatitude weather predictability (e.g., Liebmann and Hartmann 1984; Weickmann et al. 1985; Blade and Hartman 1995; Jones et al. 2004)
- An important source of predictability on the subseasonal scales.
- As a multi-scale process, it provides an ideal test bed to evaluate model physics across different spatio-temporal scales.
- MJO is stronger in boreal winter than in boreal summer.

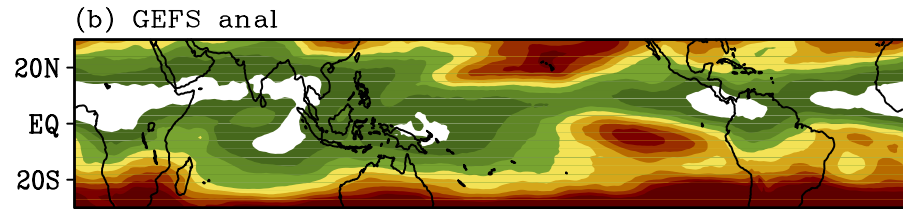
Variance of band-pass filtered U200

20–100 day U200 variance (May to Oct, 85–12)

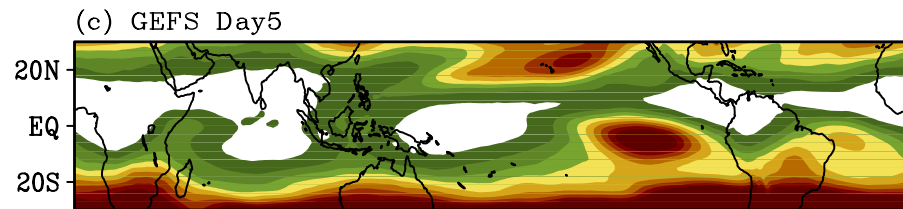
ERA-Interim



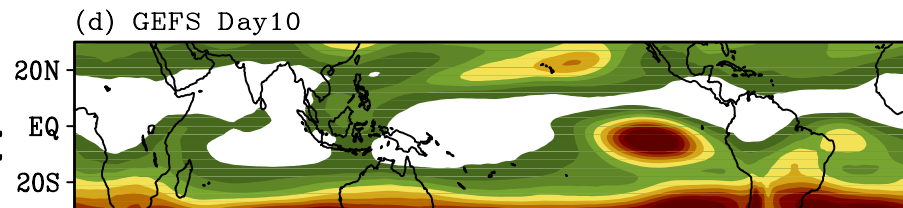
GEFS analysis



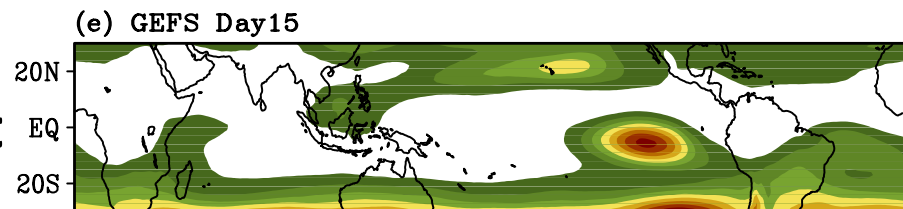
GEFS 5-day fcst



GEFS 10-day fcst

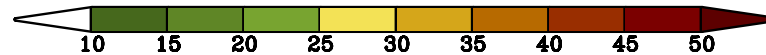


GEFS 15-day fcst



0 60E 120E 180 120W 60W 0

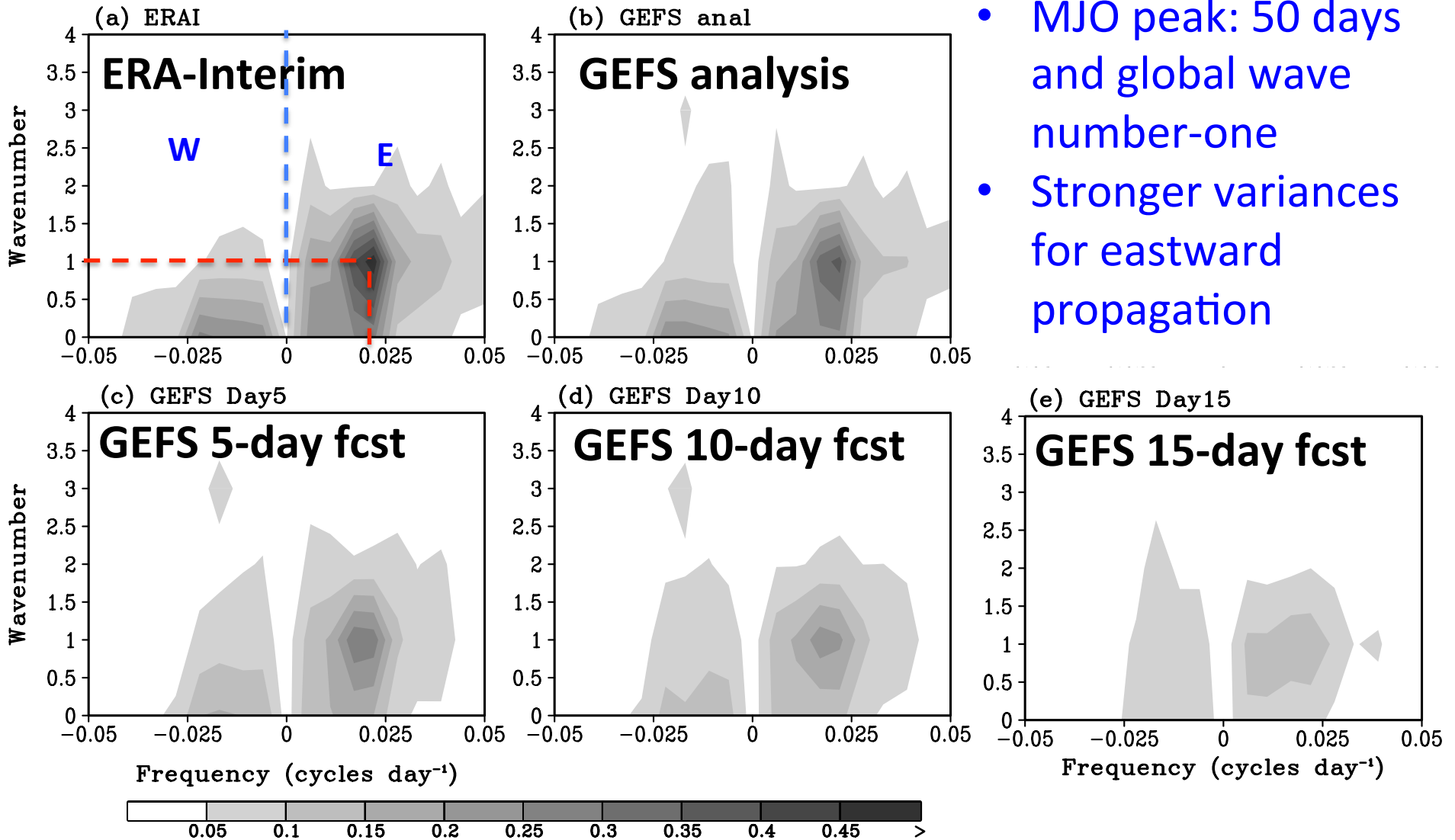
Unit : [m s⁻¹]²



10 15 20 25 30 35 40 45 50

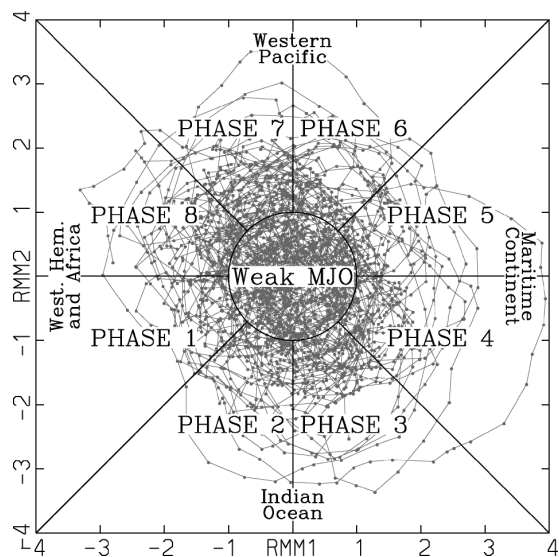
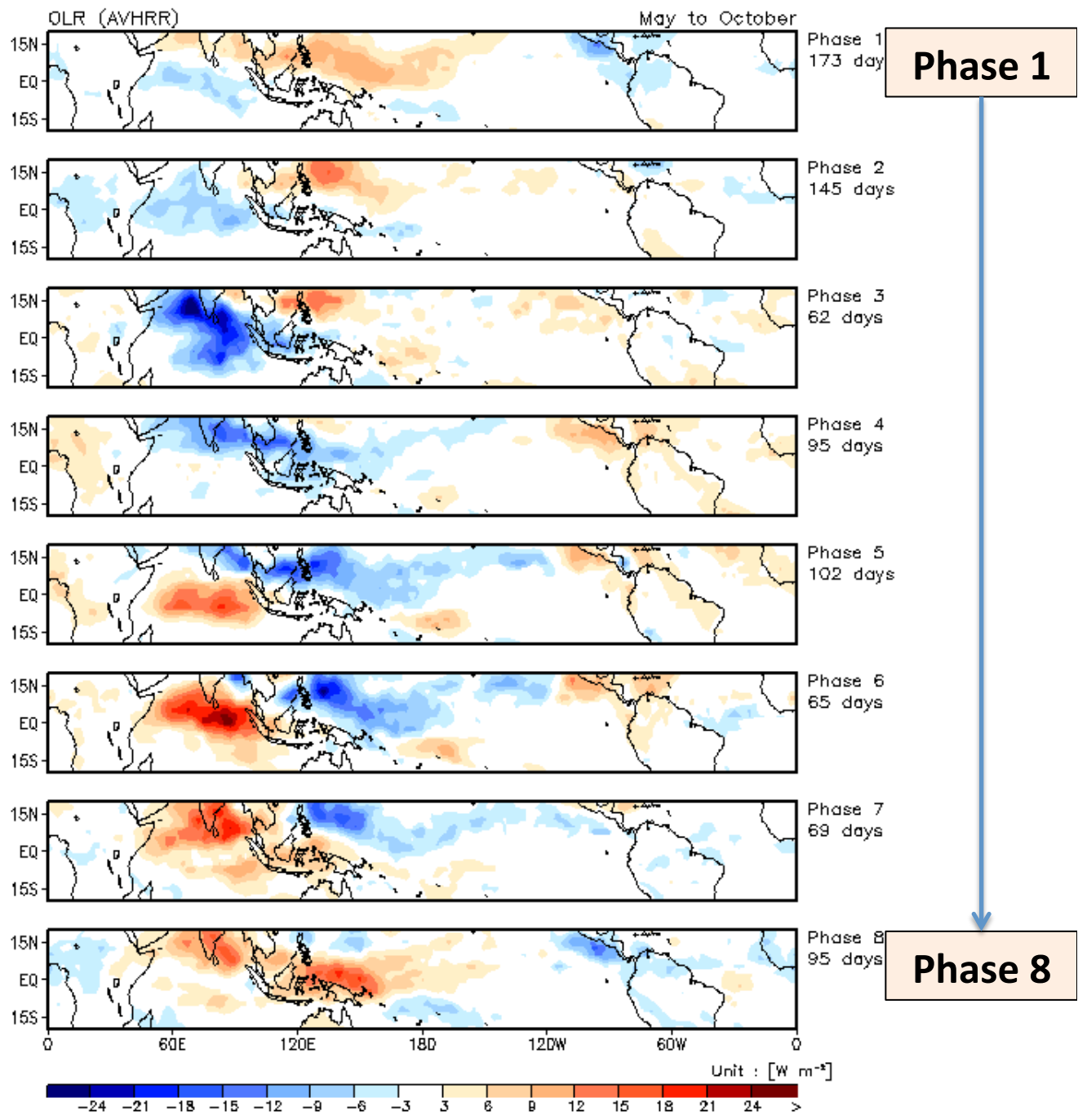
Wavenumber-frequency Diagram

Equatorial U200 Space-Time Spectra (May–Oct, 1985–2012)



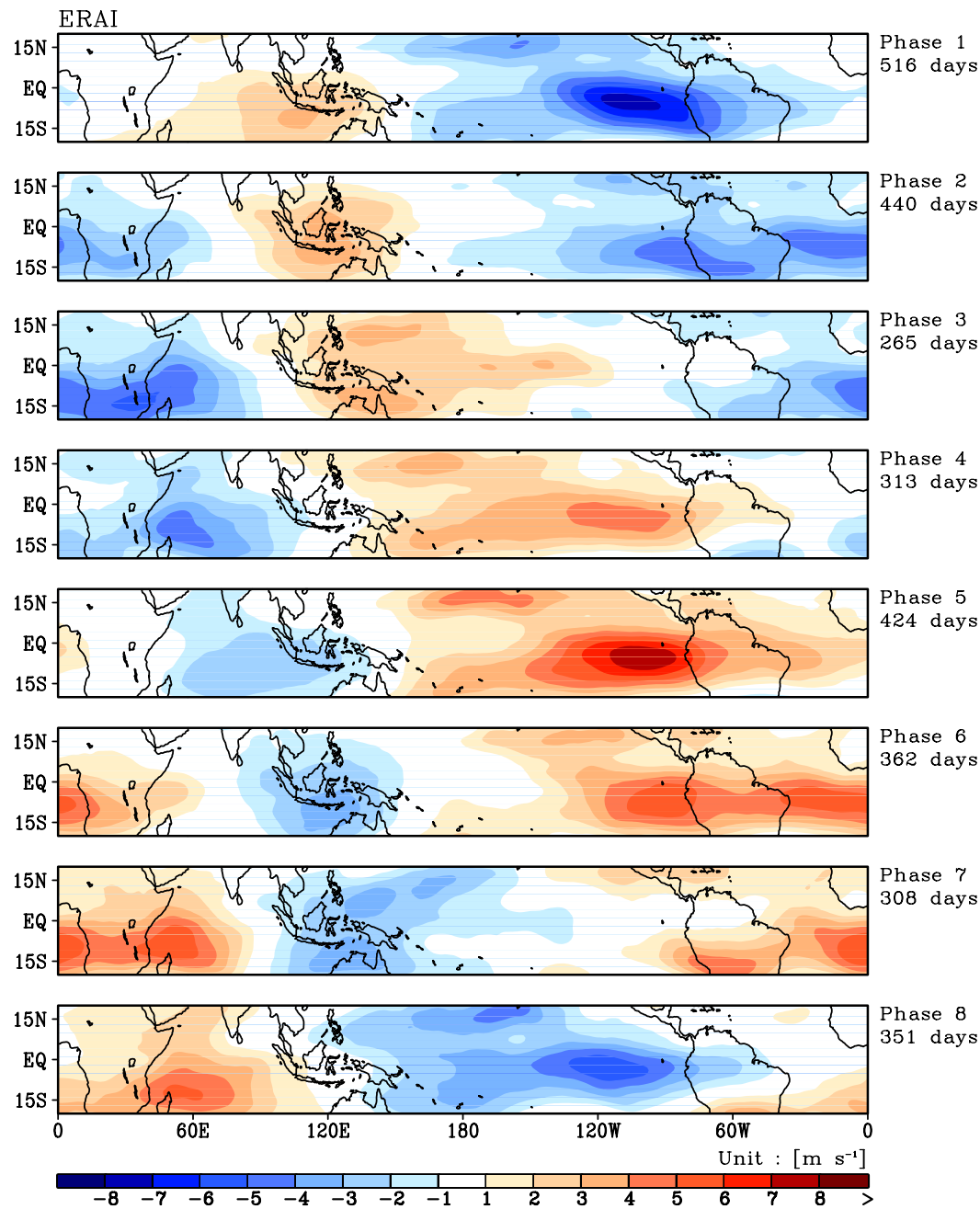
Composites based on the Real-time Multivariate MJO (RMM) Index: NOAA OLR

MJO Life cycle composite



Wheeler and Hendon 2004

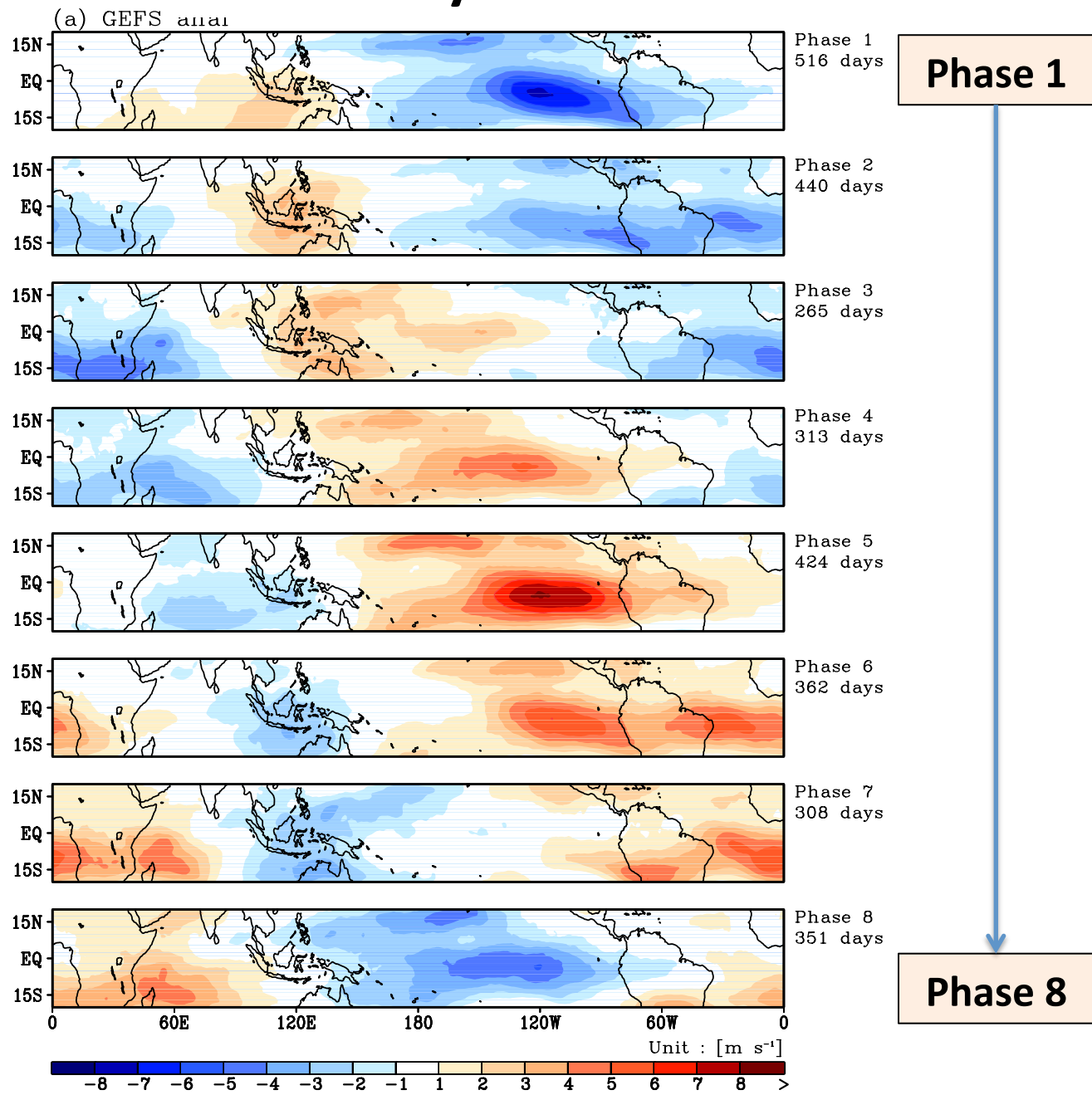
MJO Life **ERA-Interim: U200** (2012)



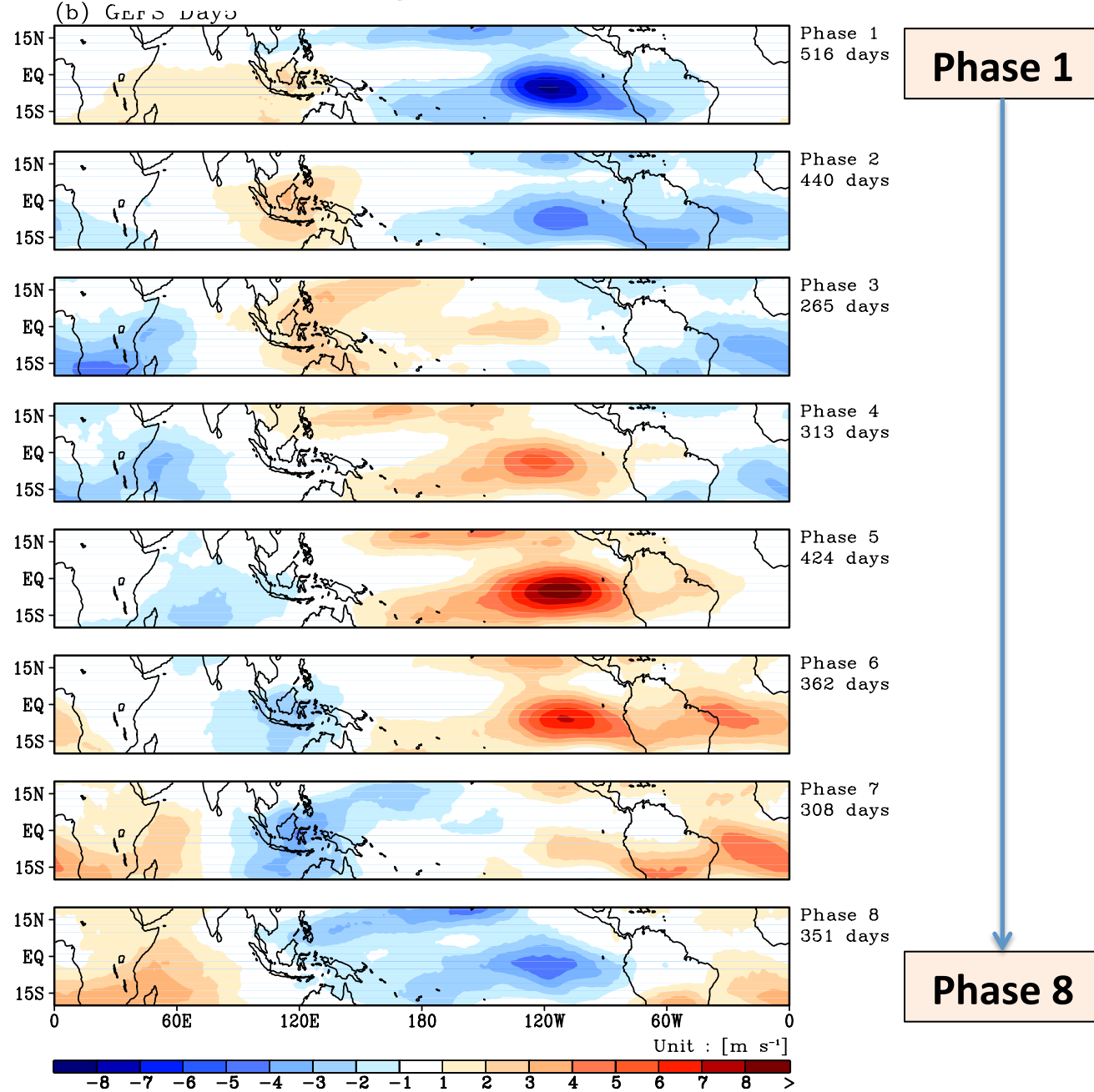
Phase 1

Phase 8

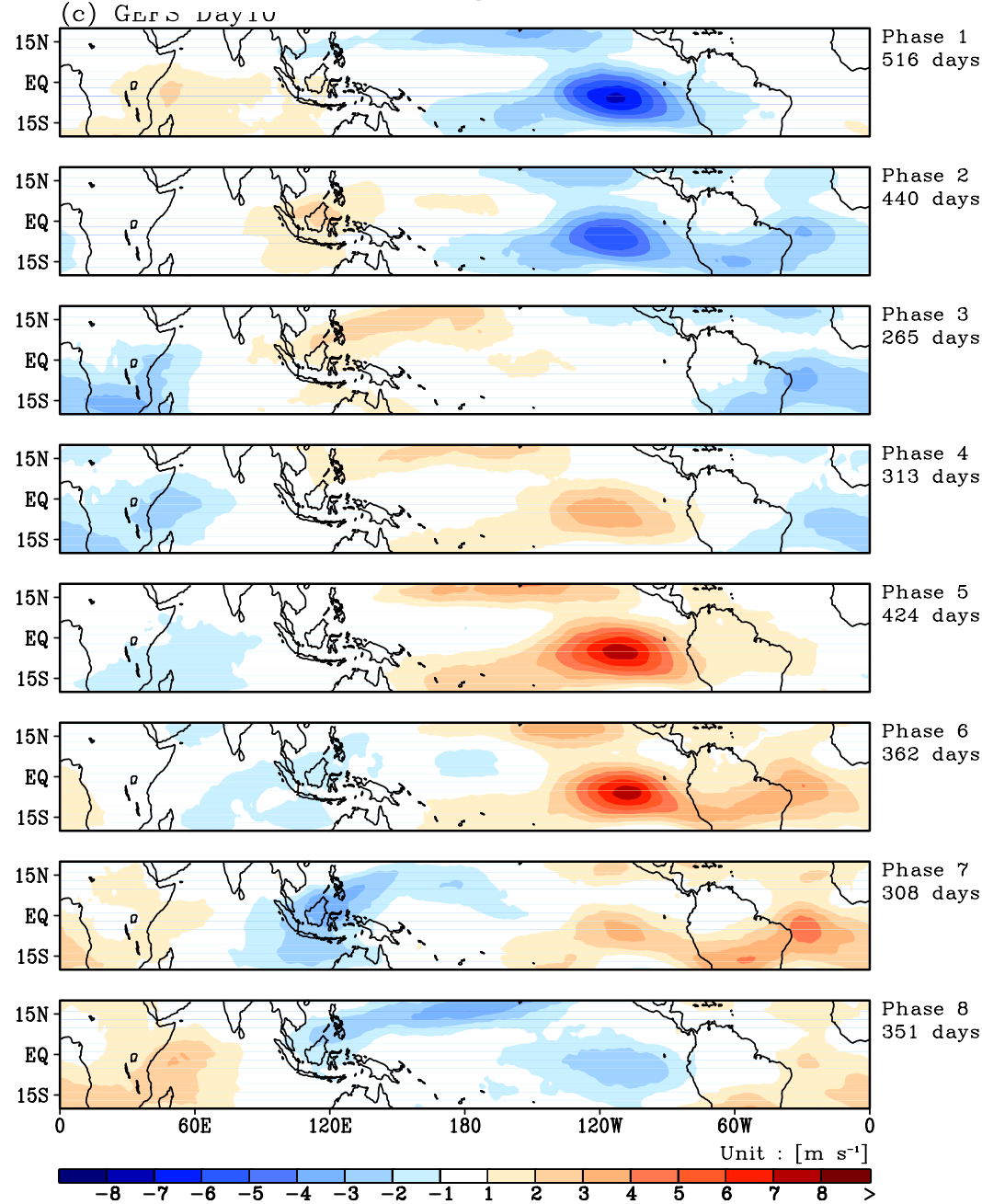
MJO Life Cycle **GEFS Analysis: U200** (12)



MJC GEFS 5-day fcst: U200 (2012)



MJO GEFS 10-day fcst: U200 ⁽²⁾

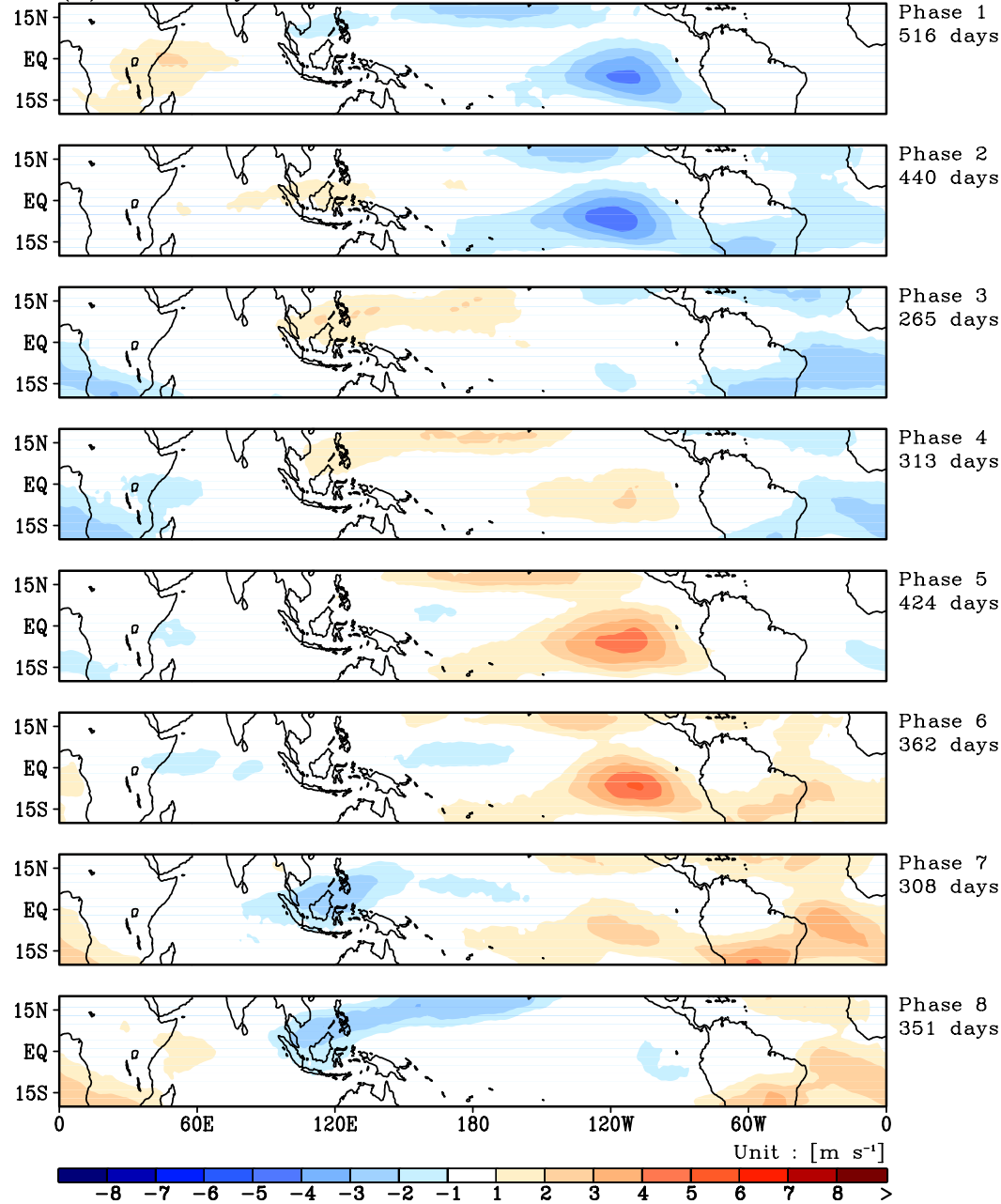


Phase 1

Phase 8

M GEFS 15-day fcst: U200 (012)

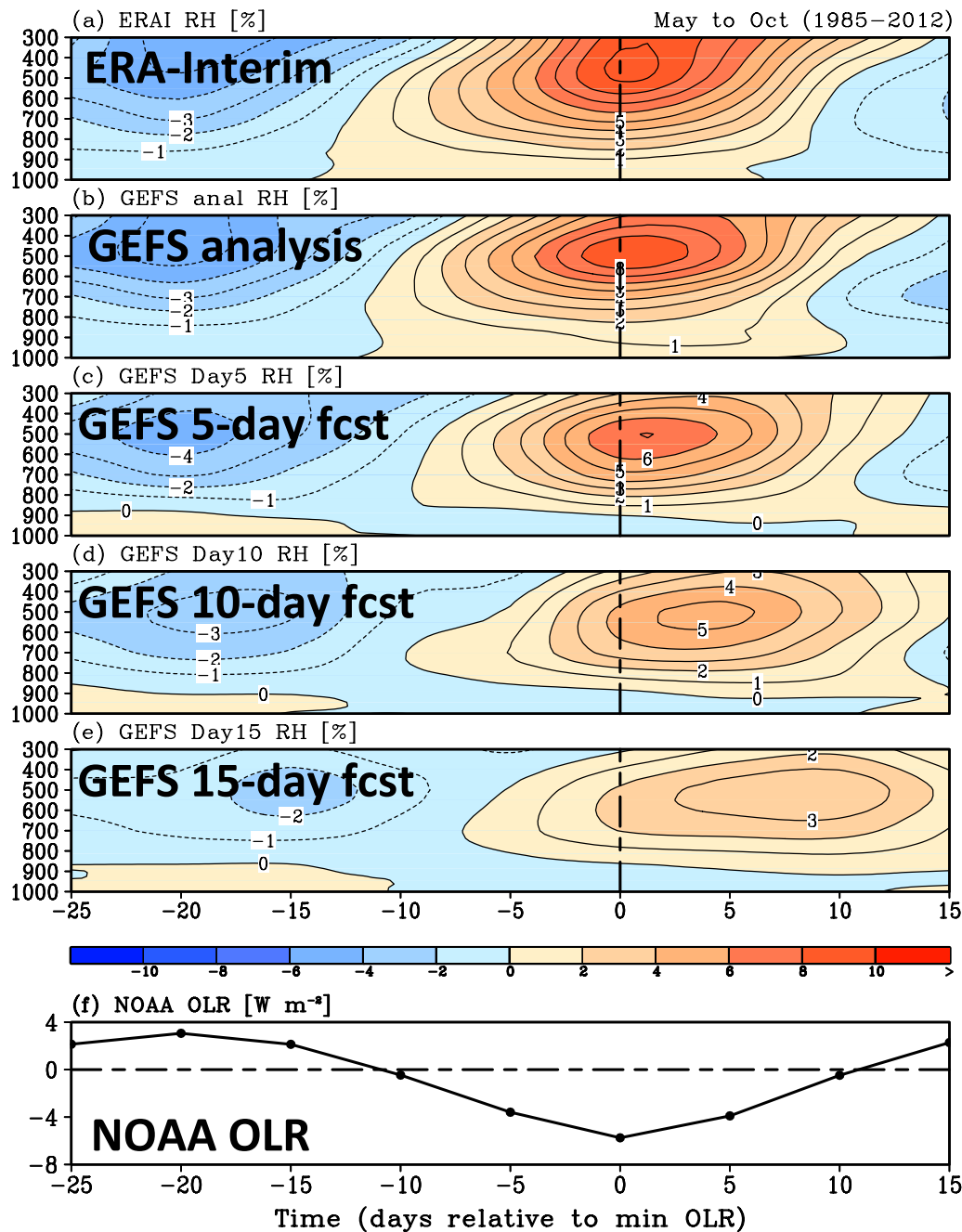
(d) GEFS Day15



Phase 1

Phase 8

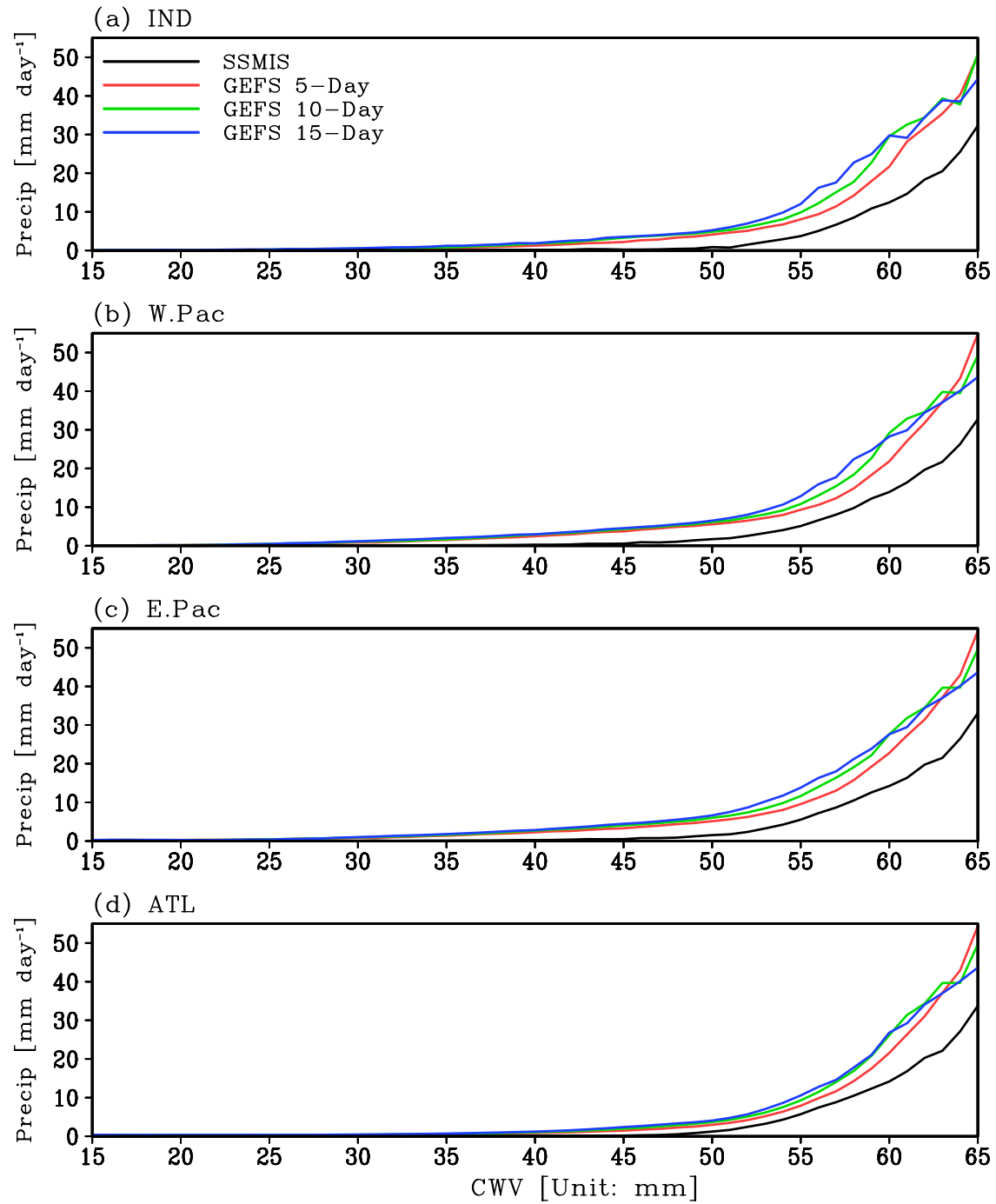
Composites by local MJO index (10S–10N,50E–100E)



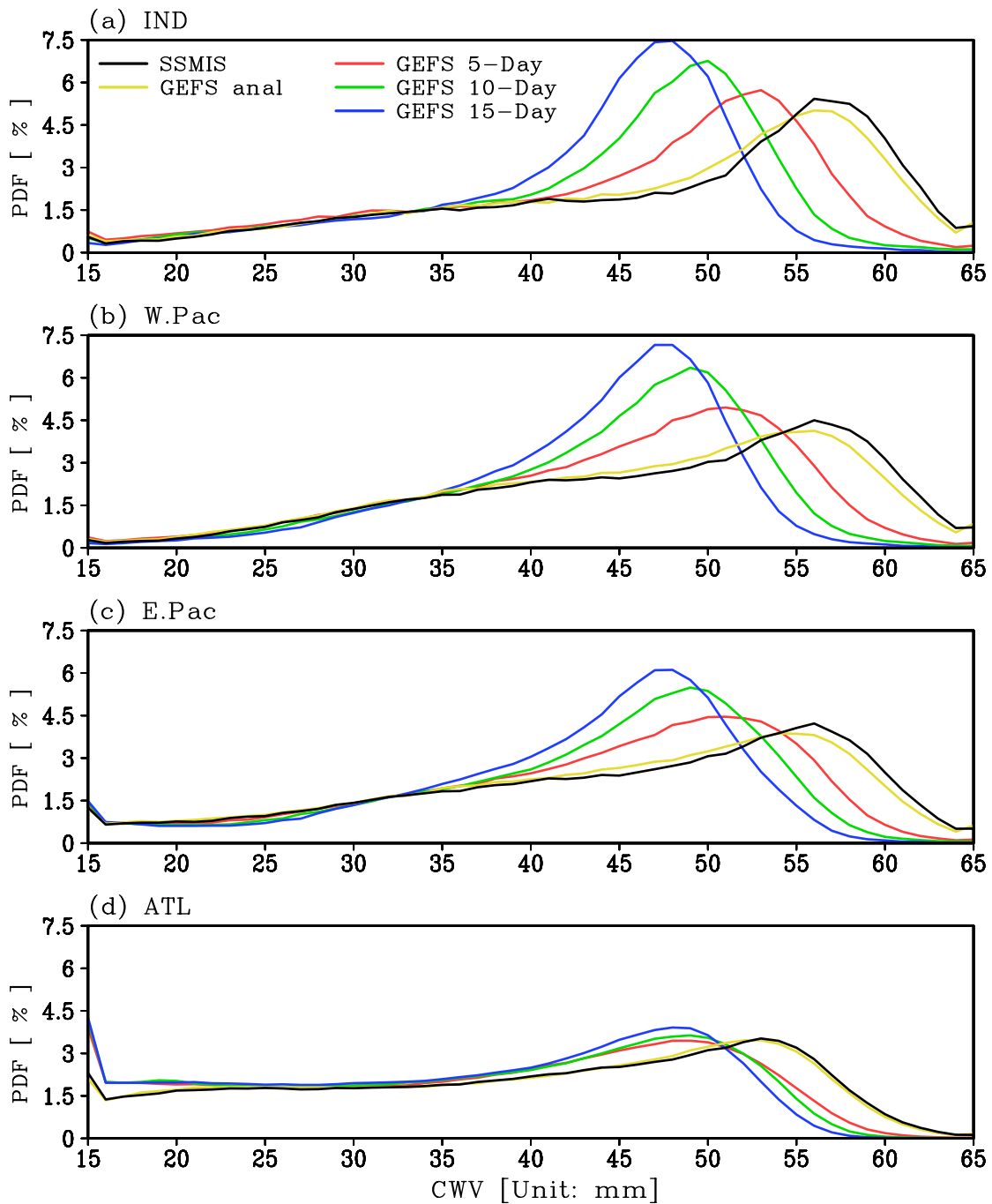
- A local MJO index was defined based on the band-pass filtered OLR.
- The magnitude of the RH anomalies weakens with the lead time.
- Are cumulus congestus and stratiform precip. under-predicted in GEFS-R?

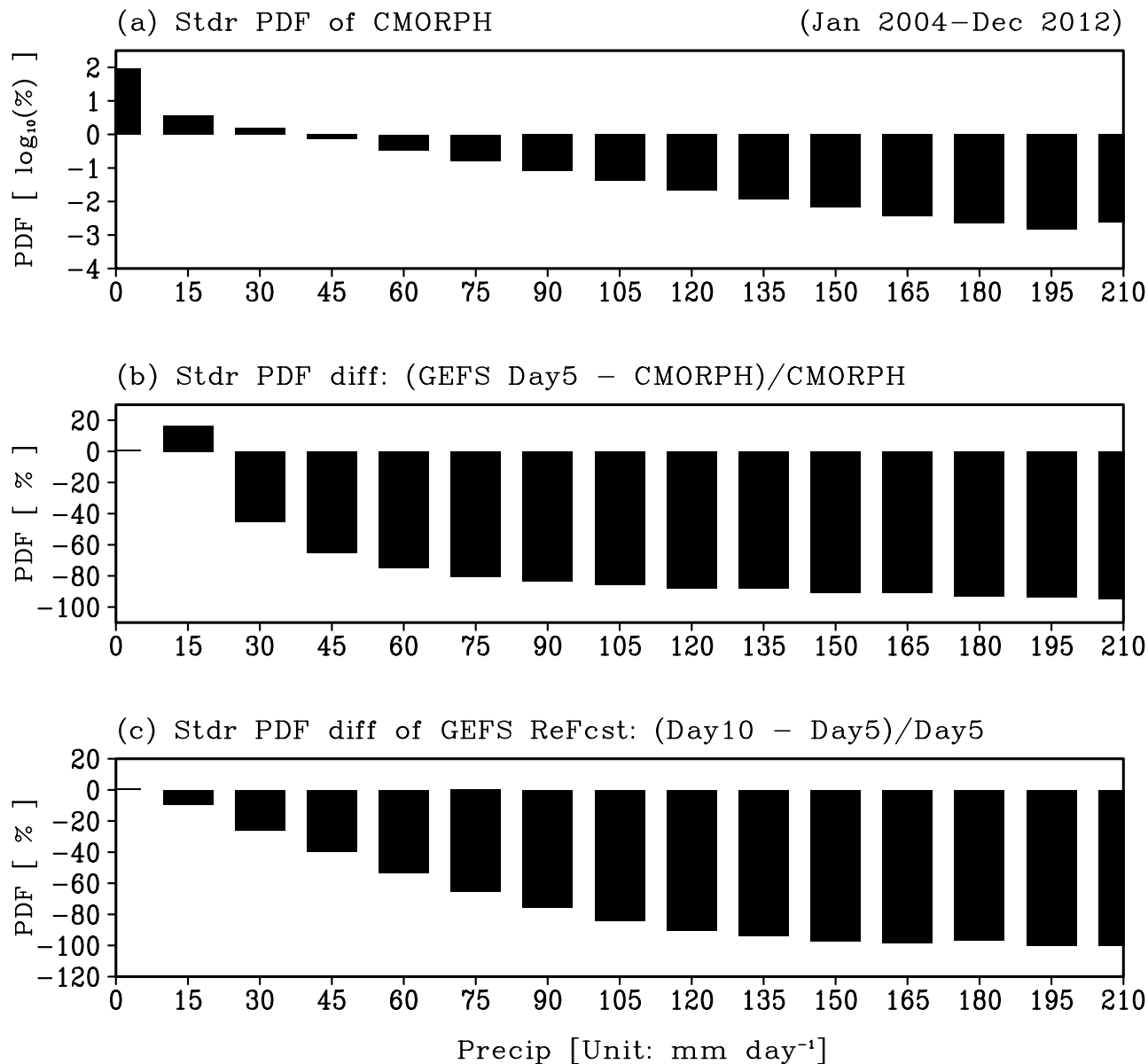
2. Precipitation Processes

Precip [mm day⁻¹] vs. CWV (2004–2012)



CWV Distribution, 2004–2012 [%]





Precipitation $\geq 30 \text{ mm day}^{-1}$ is under-predicted in GEFS-R
 The negative biases increase with the forecast lead time.

Summary of Preliminary Analyses

- GEFS captures the seasonality of TC activity reasonably well, but large errors are found in the spatial distribution of TC genesis.
- Biases in TC genesis can be attributed to the mean state errors over the West and East Pacific, such as the weaker monsoon trough and subtropical ridge, and the southward displaced ITCZ.
- High FAR near the West African coast can be attributed to the hyperactive convection and the resultant strong AEWs in GEFS.
- **Better prediction of the large-scale circulation and synoptic-scale waves can help improve TC forecasts.**
- MJO signals weaken quickly with increasing forecast lead time; the pre-moistening of the lower and middle troposphere was not well captured by the GEFS reforecasts.
- Significant dry bias in CWV was found in GEFS-R, and moderate to heavy precipitation was under-predicted in GEFS-R.

Work Plan: Physics-based Evaluations

1. Prominent Motion Systems

– Blocking

- Frequency of occurrence; extreme weather events; predictability

– Teleconnection

- Remote impacts of low-frequency perturbations

2. Specific Physical Processes

– Q1/Q2 diagnoses

- Shallow and deep heating modes; distribution and impacts

– Model representation of different cloud regimes

- Evaluated against satellite data; employ satellite simulators

– Sensitivity tests on model physics and resolution

Two Phases

- Phase 1: The products will be developed and tested using the GEF5 reforecasts.
- Phase 2:
 - Generalize the products for other models;
 - Carry out an inter-model comparison using the HIWPP data.

END OF PRESENTATION